

Milford Rodgers Bedrock Compilation Sheet 2 (paper)

Map

NOTICE !

Bedrock quadrangle 1:24,000 scale compilation sheets for the Bedrock Geological Map of Connecticut, John Rodgers, 1985, Connecticut Geological and Natural History Survey, Department of Environmental Protection, Hartford, Connecticut, in Cooperation with the U.S. Geological Survey, 1:125,000 scale, 2 sheets. [minimum 116 paper quad compilations with mylar overlays constituting the master file set for geologic lines and units compiled to the State map, some quads have multiple sheets depicting iterations of mapping]. Compilations drafted by Nancy Davis, Craig Dietsch, and Nat Gibbons under the direction of John Rodgers.

Geologic unit designation table translates earlier map unit nomenclature to the units ultimately used in the State publication.

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31 May 17
 Structure
 U/D

DEPARTMENT OF THE INTERIOR
 UNITED STATES GEOLOGICAL SURVEY

PREPARED IN COOPERATION WITH
 THE STATE OF CONNECTICUT
 GEOLOGICAL AND NATURAL HISTORY SURVEY

GEOLOGIC QUADRANGLE MAP
 BEDROCK GEOLOGY
 MILFORD QUADRANGLE, CONNECTICUT
 GQ-427

EXPLANATION

In the descriptions of rocks below, minerals are listed in approximate order of decreasing abundance, unless otherwise noted. Grain sizes are as follows: fine, less than 1 mm; medium, 1-5 mm; coarse, more than 5 mm. Descriptions of rocks do not include randomly distributed minerals formed by retrograde metamorphism after the climax of the latest progressive regional metamorphism. The retrograde metamorphism resulted in slight to complete alteration of mineral assemblages: kyanite and staurolite mainly to sericite, biotite to chlorite, garnet to sericite and chlorite, pyrite to hematite, and rutile and ilmenite to leucosane.

UNCONFORMITY
 Allington Metadiabase
 Oa, mainly medium- to fine-grained, dark greenish-gray to dark-gray, intrusive metadiabase in chlorite zone (east of garnet isograd) and equivalent medium- to fine-grained, gray-green to dark greenish-gray amphibolite in zones of higher grade metamorphism. Porphyritic in many places; numerous microcline and perthite inclusions. Metadiabase is composed of actinolite hornblende, altered plagioclase, epidote, chlorite, muscovite (mainly sericitic), and accessory magnetite or ilmenite, sphene, apatite, and calcite. Amphibolite is composed mainly of hornblende, sodic andesine, epidote, chlorite, and abundant accessory sphene. Boundaries general. Near Connecticut. Porphyry contains numerous inclusions of Oronoque Member of Derby Hill Schist 1-5 feet thick inter-layered with sill-like and gently inclined dike-like bodies of metadiabase 1-5 feet thick. Not known to intrude Wepawaug Schist.

Malby Lakes Volcanics
 Omj, medium- to fine-grained, medium light-gray impure crystalline limestone approximately 12 feet thick, composed of calcite, quartz, chlorite, muscovite, and minor biotite, sphene, apatite, clinzoisite, and tourmaline.
 Oms, medium- to fine-grained, thinly layered metasedimentary rocks similar to Derby Hill Schist.
 Omp, pyroclastic schist, meta-agglomerate or meta-tuff characterized by aluminous greenish-gray to dark greenish-gray matrix with angular fragments as much as 1 inch thick and 5 inches long rather evenly distributed in a medium- to fine-grained matrix of light to dark greenish-gray schist composed of quartz, plagioclase, quartz, muscovite, chlorite, and minor garnet, biotite, magnetite or ilmenite, tourmaline, rutile, apatite, and pyrite. Black hornblende prisms rather randomly arranged on foliation planes in many places. Moderate greenish-yellow lenses, and knots from less than 1 mm to more than 1 inch thick characteristic regardless of metamorphic grade; composed largely of epidote and quartz.
 Omi, medium- to fine-grained, well-sorted, medium light-gray impure crystalline limestone approximately 12 feet thick, composed of calcite, quartz, chlorite, muscovite, and minor biotite, sphene, apatite, clinzoisite, and tourmaline.
 Omo, medium- to fine-grained, thin layered metasedimentary rocks similar to Derby Hill Schist.
 Ood, Oronoque Member, mainly slabby to thinly laminated, medium- to fine-grained, greenish-gray to medium dark-gray quartz-rich and albite paragneiss with schistose to phyllitic partings and layers containing abundant muscovite and chlorite; in some places in western part of area contains green or brown biotite and a little garnet; at least one schistose band north of Devon also contains staurolite and kyanite, but these minerals not characteristic; accessory minerals are ilmenite, magnetite, rutile, sphene, apatite, tourmaline, zircon, pyrite, and dust-like carbon. Quartz-rich layers from 1 mm to several feet thick are much more abundant than in underlying unit, especially in western half of quadrangle. Upper part of member also contains minor ungrouped epidote-rich greenish-gray and amphibolite similar to those characteristic of undivided Malby Lakes Volcanics.
 Odl, fine-grained, light- to medium-gray crystalline limestone containing streaks, lenses, pods, and bands of moderate yellowish-green to dusky yellow-green serpentine and abundant accessory magnetite.
 Ods, amphibolite mineralogically similar to that characteristic of Malby Lakes Volcanics in staurolite and kyanite zone.

Southington Mountain Schist
 Mainly interlayered medium- to fine-grained silvery-gray to medium-gray muscovite schist and slabby medium light-gray to dark-gray paragneiss. Schist composed of quartz, muscovite, chlorite, sodic plagioclase, and accessory ilmenite, rutile, tourmaline, zircon, and dust-like carbon. Paragneiss composed of quartz, muscovite, biotite, garnet, albitic plagioclase, hornblende, chlorite, and accessory sphene, apatite, and zircon. Poorly exposed in this area but characterized by ribbonlike banding in type area in Southington quadrangle.

Outcrop or outcrop area
 Generalized in many places, and may include some places where bedrock merely close to surface. Shows distribution of most but not all outcrops in map area.

CONTACT
 Dashed where approximately inferred; short dashed where indefinite or inferred; quiet where location, extent, or nature of contact uncertain.

INFERRED SYNECLINE
 Showing trace of axial plane, and bearing and plunge of axis.

PLANAR AND LINEAR FEATURES IN METAMORPHIC ROCKS
 Symbols may be combined.

Vertical
 Foliation or flow cleavage
 Formed by parallel alignment of minerals, such as mica in schist or gneiss. Generally about parallel to bedding or relic bedding in banded metasedimentary rocks, although foliation crosses contacts of pyroclastic schist, Omp, and some beds between Malby Lakes Volcanics and adjacent formations west of Wepawaug River where rocks have been tightly folded. Minor microcline measured dip, quartz where uncertainty of dip inferred. Symbol for inclined foliation with number indicates direction of dip inferred. Symbol accompanied by ± indicates dip locally greater or less than that recorded.

Inclined
 Fracture cleavage
 Formed by parallel or nearly parallel arrangement of axial planes of symmetrical crenulations or by alignment of long limbs of asymmetrical crenulations in schist. Number indicates measured dip.

Horizontal
 Crumpled or contorted foliation and (or) banding
 Direction of long line indicates generalized strike.

Horizontal
 Lineation
 Mineral lineation unless otherwise noted; C, crenulation axis; FA, fold axis.

Isograd
 Denotes approximate eastern limit of zone in which mineral named (kyanite, staurolite, biotite, or garnet) in highest grade macroscopic index mineral in metamorphosed pelitic rocks. Position based in part on known positions of isograds in Wepawaug Schist in adjacent Ansonia quadrangle. Minor microcline measured dip, quartz where uncertainty of dip inferred. Symbol for inclined isograd in the northern corner of the map area, especially near the Allington Metadiabase, but generally not recognized in the field.

Macroscopic-mineral localities
 K, kyanite; S, staurolite.

Abandoned quarry
 Type locality.

NOTES ON GEOLOGY

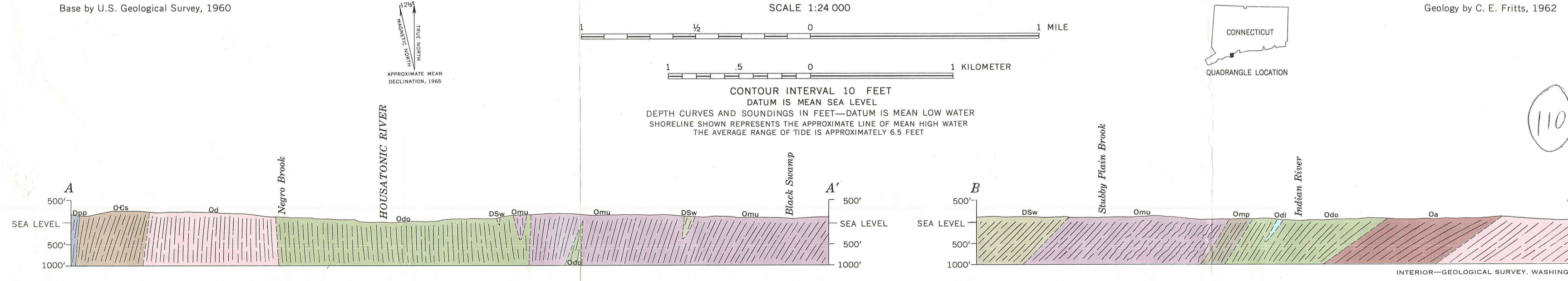
Stratigraphy, structure, and regional metamorphism. The Southington Mountain and Derby Hill Schists, the Malby Lakes Volcanics, and the Wepawaug Schist form part of a sequence of metasedimentary and metavolcanic rocks of Paleozoic age, which in Middle to Late Devonian time was tightly folded, intruded by the Prospect Gneiss, and subjected to progressive regional metamorphism ranging from chlorite to kyanite grade (Fritts, 1962a, 1962b, in press). The principal structure formed in this area at that time is a complex major syncline, the trough of which is occupied by the Wepawaug Schist. This fold, known as the Wepawaug syncline, plunges north-northeastward and also underlies parts of the Woodmont, Ansonia, New Haven, Naugatuck, and Mount Carmel quadrangles. The disappearance of the Malby Lakes Volcanics along the western limb of the Wepawaug syncline in the northern part of the Milford quadrangle is interpreted as evidence that this formation had an eastward or southeastward dip at the time the Wepawaug Schist was deposited unconformably above it. Although the intrusive Allington Metadiabase also appears to be confined mainly to the eastern limb of the Wepawaug syncline, some of the small bodies of amphibolite in the Derby Hill Schist on the western limb of the syncline may be related to the metadiabase.

A gradual westward increase in metamorphic grade of the rocks in this area coincides with a westward increase in the complexity of geologic structure. In the eastern part of the quadrangle, the metamorphic grade of the rocks is low and geologic structure is rather simple. Phyllitic schists, greenish, and low-grade metadiabase predominate. In general, foliation in metasedimentary rocks there is parallel or nearly parallel to bedding. However, small tight folds are outcrops of even the least metamorphosed rocks at Morning-side, and field relations in the Ansonia quadrangle, in part, indicate that foliation is not necessarily parallel to the boundaries of metavolcanic rocks such as the pyroclastic schist of the Malby Lakes Volcanics. West of Milford, the metamorphic grade of the rocks is higher, and mica schists, paragneisses, and amphibolites predominate. Geologic structure there also is more complex. Tongues of Wepawaug Schist, for example, projecting southwestward into the Malby Lakes Volcanics apparently are in the troughs of minor nearly isoclinal synclines in a zone of intense folding near or west of the axial plane of the main (Wepawaug) syncline. Foliation obviously crosses stratigraphic boundaries near the axial planes of these folds. The author believes that the large tongue of Malby Lakes Volcanics projecting southwestward along to Long Island Sound at Laurel also represents a tight, north-plunging syncline along or near the axial plane of the Wepawaug syncline. This interpretation was facilitated by the recent discovery of seven outcrops of the Malby Lakes Volcanics near Fort Trumbull and along the Wepawaug River at Milford by John Rodgers and J. E. Sanders (John Rodgers, written commun. 1964). The location of the contact between the Wepawaug Schist and the Malby Lakes Volcanics east and south of Baldwin Swamp, however, is uncertain because of a lack of outcrops. Thus it is possible that the Wepawaug Schist in the trough of this fold projects as far southward as the U.S. Military Reservation near Meadowside School.

There is no indisputable evidence of large-scale faulting in this area, but it is possible that faults of Triassic age mapped in adjacent quadrangles extend into this one. The Mixville fault, for example, which forms the western boundary of the Newark Group of Late Triassic age in the Southington and Mount Carmel quadrangles, apparently extends south-southwestward into the Ansonia quadrangle and may continue into the Milford quadrangle, perhaps somewhere in the drift-filled valley of the Wepawaug River. The fault has not been inferred here, however, because geologic structure in this area can be explained without it. The throw on the Mixville fault is not known to be more than a few hundred feet, and a horizontal component of displacement has not been proved. Furthermore, all faults must hinge out somewhere, and there is no conclusive evidence that this fault extends south of the town of Orange in the Ansonia quadrangle. On the other hand, it is more likely that the Great Fault, which forms the eastern boundary of the Newark Group in central Connecticut and swings southwestward toward Morris Cove in the New Haven quadrangle, continues southwestward into the Milford quadrangle somewhere beneath Long Island Sound.

REFERENCES CITED

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BEDROCK GEOLOGIC MAP OF THE MILFORD QUADRANGLE, FAIRFIELD AND NEW HAVEN COUNTIES, CONNECTICUT

By
 Crawford E. Fritts
 1965

65
 Strike + dip of foliation
 plunge of lineation
 110
 Normal fault
 V - with room side
 D - down the room side
 J - Reverse or thrust fault
 T - (E Devonian)
 T-O with mud side